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Full Length Research Paper

Hydrogeochemical characterisation of Shallow and Deep aquifers in and around Limestone Mining Area from Gadchandur, Chandrapur District (Maharashtra), India

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ABSTRACT

Evaluation of ground water quality and hydro-geochemistry of shallow and deep aquifer (with respect to guideline values of WHO 2004, geochemical classifications, ion exchange processes and mechanisms controlling ground water chemistry) and evaluation of water quality (with respect to different sodium hazards, Ca²⁺/Mg²⁺ ratio) for irrigation purpose in mining area were the aims of the present study. Accordingly ground water samples were collected from shallow, deep aquifers of Gadchandur, Chandrapur in pre and post monsoon seasons 2010. Quality of the water samples collected from the deep aquifer found satisfactory in comparison with water samples collected from the shallow aquifer. Geochemical nature of shallow aquifer was earth alkaline with increased portion of alkalis with prevailing bicarbonate, followed by chloroalkaline disequilibrium type of ion exchange process. Chemical weathering of rock forming minerals was the major driving force controlling shallow aquifer water chemistry. Hydro-geochemistry of deep aquifer in pre-monsoon season hydrogeochemistry of the deep aguifer shifted towards alkaline with sulphate and bicarbonate. Ion exchange process in this aguifer showed complete dominance of base exchange reaction in both the seasons. Chemical weathering along with evaporation was the two major driving forces controlling the water chemistry of the shallow and deep aquifer in pre-monsoon and post-monsoon seasons. RSC (residual sodium carbonate), %Na, SAR (sodium adsorption ratio), Ca²⁺/Mg²⁺ ratio and concentration of CI, F in irrigation water revealed that some water samples were not good enough for irrigation in pre-monsoon season in comparison to that of post monsoon season.

Keywords: Shallow aquifer, Deep aquifer, Hydro-geochemistry, Irrigation water

INTRODUCTION

The ever-growing utilization of water resources have led to lowering of water table as well as degradation of its quality in the past few decades. Number of studies on groundwater and surface water quality in parts of India with respect to drinking and irrigation purposes have been carried out extensively by Central Ground Water Board, Geological Survey of India as embodied in various publications. Mining activity threatens the quality and



Figure1. The location of study area and water sample points

quantity of surface and ground water resources in many parts of world (Allen et al., 1996; Choubey, 1991; Gupta, 1999; Khan et al. 2005; Singh et al. 2008; Singh 1998; Tiwary 2001; Soni, 2007). Mining activity, thus in many areas can seriously pollute surface and groundwater resources. The various studies on interrelationship between water quality, geology and mining activities have been carried out (Singh et al., 2010). The present study was performed with an objective to evaluate ground water quality and hydro-geochemistry of shallow and deep aquifer water quality for irrigation purpose in and around limestone mining area from Gadchandur, Chandrapur District Maharashtra.

STUDY AREA

The study area is located in south-western part of Chandrapur District (around Gadchandur), Maharashtra state and lies between latitudes 19 38'30" to 19 50'30"N and longitudes 79 04'00" to 79 11'00"E (Figure 1). The area under investigation is covered under Survey of India topographic sheet No. 56 M/1, M/2. The study area is around 90 Km away from Chandrapur and it is approachable by Rajura - Korpana Road.

In the study area dendritic (major), trellis and combination of dendritic and (minor) parallel drainage patterns have been observed. The climate is semiarid with minimum temperatures drops up to 10°C in winter (January) and 48°C in summer (May). The study area receives an average annual rainfall 1132.21 mm.

The study area covers Geological formations *viz*. Limestone-shale and Limestone sequences belonging to Penganga Group of Godavari valley, Lower Gondwana sediments i.e. Talchir, Barakar and Kamthi formations, Deccan Trap Basalts, Alluvium, soil and Laterite.

The two mining areas fall in the study area *viz.*, i) Naokari Limestone Mining Area in northern part and ii) Manikgarh limestone mining area in the southern part. Naokari Limestone Mining area occupied by limestone and shale sequences belonging to Penganga Group and the rocks of Gondwana Supergroup in the extreme northern part. Manikgarh limestone mining area in the southern part study area is occupied by Penganga Limestone, along with Deccan Trap basaltic flows with laterite capping. In the study area three types of soil textures are found Clayey Soil 43.08%, Sandy Soil 13.96 % and Gravelly Soil 42.97 %.

The groundwater occurs under phreatic conditions in all the water bearing formations, in weathered fractured zone (limestone), fractured and vesicular basalts (shallow aquifer) and under semi-confined to confined conditions in the deeper aquifers. The thickness of weathered zone varies from 5m to 18m in Penganga rocks and 2 to 12 m in basaltic formations. In Penganga rocks groundwater occurs in fractures and cavities in krastified limestones and joints of sandstone. In Gondwanas, groundwater occurs in the pore spaces of Barakars and Kamthi sandstones and in deeper aquifers it occurs under semiconfined conditions. In alluvium, groundwater exists in sandy layers under phreatic conditions. Local confined conditions are sometimes simulated by clay lenses. The water bearing zone occurs at a depth of 1.8 to 13.5 m in shallow aquifers in the present area.

MATERIALS AND METHOD

Sixty two (31×2) water samples (17 shallow aquifer water samples, 14 deep aquifer water samples) were collected from the study area in pre monsoon 2010 and post monsoon 2010. The sampling stations are shown in Figure 1.

All the analytical methodologies including sample collection, had been carried out as per Standard Methods for the Examination of Water published by APHA (1998). Physical parameters like pH by pH meter and electrical conductivity (EC) were determined by portable conductivity meter. Total dissolved solids (TDS) were determined by gravimetric method. Chemical parameters like sulphate (SO₄) was determined by UV spectrophotometric model UV- 1201A, whereas, total hardness (TH), TA measured by the titration of the sample with standard N/50 H₂SO₄. Ca hardness (CaH), calcium (Ca²⁺) and magnesium (Mg²⁺) were determined by standard EDTA titrimetric method. Carbonate (CO_3^2) and bicarbonate (HCO₃) were estimated by determined by titration of the sample with standard N/50 H_2SO_4 , using mixed indicator consisting of methyl red and bromocresol green (1:1). Chloride (Cl) was determined by argentometric titration method. Sodium (Na⁺) and potassium (K^{+}) had their measurement by the techniques of emission Flame Photometry. Fluoride (F) was measured by spectrophotometrically by using Simadzu make UV spectrophotometer (model UV-1201A) with automatic sampler (model ASC-5). Rest of the chemical parameters like total iron (Fe), nitrate (NO₃) were measured by using Simadzu make UV spectrophotometric model UV-1201A.

RESULTS AND DISCUSSION

Water quality of shallow and deep aquifer samples

Statistical summary of analytical data of the water samples collected from shallow and deep aquifer are described in Table 1 and Table 2.

Table 3 describes that, 100% of the water samples collected from shallow aquifer were found within safe limit as per with the provisional guideline values of WHO 2004 with respect to pH and SO_4 point of view. pH, Na and SO_4 of all the water samples collected from deep aquifer in

| Parameters | Range (Pre- monsoon 2010) | Average (Pre-monsoon 2010) | SD | Range (Post monsoon 2010) | Average (Post-monsoon 2010) | SD |
|------------|------------------------------|-------------------------------|--------|------------------------------|--------------------------------|--------|
| pН | 7.6-8.3 | 7.9 | 0.19 | 6.9-8.5 | 7.52 | 0.52 |
| EC | 463-3410 | 1229.82 | 955.31 | 664-2800 | 1436.57 | 963.39 |
| TDS | 296-2114 | 785.47 | 599.54 | 425-1792 | 924.5 | 619.42 |
| TH | 232-760 | 414.88 | 170.85 | 304-1136 | 569.21 | 311.34 |
| Са | 35-163 | 69.70 | 28.30 | 43-291 | 120.14 | 83.46 |
| Mg | 19-119 | 54.23 | 23.77 | 35-155 | 69 | 35.34 |
| Na | 15-216 | 58.94 | 65.40 | 16-255 | 76.92 | 50.14 |
| К | 1-101 | 10.29 | 26.28 | 1-129 | 20.21 | 42.73 |
| CI | 28-438 | 140.17 | 145.25 | 28-316 | 152.71 | 209.09 |
| NO3 | 1-57 | 16.23 | 18.68 | 31-421 | 167.92 | 121.91 |
| CO3 | 0-0 | 0 | 0 | 0-64 | 10.28 | 21.30 |
| HCO3 | 160-548 | 309.41 | 86.83 | 160-560 | 330.57 | 80.09 |
| SO4 | 1-172 | 45.23 | 44.601 | 15-204 | 56.14 | 38.79 |

Table 1. Descriptive statistical summary of the analytical data (Shallow aquifers)

All unites have their measurement in mg/l except, 'T' is expressed in °C, 'EC' is expressed in µScm, 'TH and CaH' are expressed in mg/l as CaCO3, SD=standard deviation, BDL =below detection level

Table 2. Descriptive statistical summary of the analytical data (Deep aquifers)

| Parameter | Range (Pre- monsoon 2010) | Average (Pre- monsoon 2010) | SD | Range (Post - monsoon 2010) | Average (Post -monsoon 2010) | SD |
|-----------|------------------------------|--------------------------------|----------------|--------------------------------|---------------------------------|--------|
| nH | 76-85 | 8 00 | 0.25 | 7 1-8 6 | 0.52 | 0.46 |
| ρΠ EC | 3/18-2230 | 967 | 0.20 440.87 | 544-4280 | 0.52 | 622.84 |
| | 206 2114 | 507 | 440.07 | 249 2720 | 610.42 | 200.06 |
| 105 | 290-2114 | 597 | 190.36 | 340-2739 | 019.42 | 399.90 |
| TH | 312-995 | 362 | 136.83 | 184-1536 | 311.34 | 232.08 |
| Ca | 184-689 | 63 | 31.46 | 11-336 | 83.46 | 64.18 |
| Mg | 11-130 | 47.78 | 19.14 | 22-167 | 35.34 | 36.46 |
| Na | 19-175 | 72.07 | 50.23 | 15-175 | 50.14 | 71.57 |
| К | 1-39 | 7.85 | 12.86 | 1-126 | 42.73 | 33.136 |
| CI | 26-243 | 77.5 | 93.95 | 30-834 | 209.09 | 91.79 |
| NO3 | 2-48 | 14.42 | 12.51 | 35-487 | 121.91 | 100.15 |
| CO3 | 0-48 | 3.42 | 13.31 | 0-56 | 21.30 | 21.24 |
| HCO3 | 160-488 | 340.21 | 93.95 | 160-508 | 80.09 | 97.77 |
| SO4 | 8-97 | 35.28 | 27.97 | 18-168 | 38.79 | 55.49 |

All unites have their measurement in mg/l except, 'T' is expressed in °C, 'EC' is expressed in µScm, 'TH and CaH' are expressed in mg/l as CaCO3, SD=standard deviation, BDL =below detection level

post-monsoon and pre-monsoon seasons were found within the safe limit. Regarding the values of TDS, TH, Na⁺, Cl⁻ and NO₃ 67, 65, 88, 73 and 55% samples respectively in shallow aquifer and TDS, TH, Cl⁻ and NO₃ 91, 64, 88 and 54% samples respectively in deep aquifer were found safe.

HYDRO-GEOCHEMISTRY OF THE AQUIFERS

The Hydro-geochemical Facies

The determination of hydro-chemical facies was carried out as per the Piper trilinear diagram (1944) using Aquachem 4.0 software to assess the nature of hydrogeochemistry of both the aquifers. The piper diagram directly deciphers the groundwater quality, and as a result it is commonly used as an effective tool to specify different water-types. In general, piper diagram reveals six types with nine combinations, according to ionic placement in the diamond field (Figure 2). Plots of analytical data (Figure 2) of the water samples collected from shallow aquifer clustered in the divisions 1,2,4 and 5 of Piper diagram, which showed that alkali earths exceed alkalis, weak acids exceeds strong acid and carbonate hardness (secondary alkalinity) exceeds 50% (Table 4) type of hydro-geochemistry.

| Physical and unconfined | WHO standards for drinking | Shallow aquifer | Deep aquifer | |
|---------------------------------------|---|-------------------------------|----------------------------------|--|
| aquifer chemical Parameters (mg/l) | water, 2004 (maximum permissible limit) | Samples within safe limit (%) | Samples within safe limit (%) | |
| pH | 6.5 - 9.5 | 100 | 100 | |
| EC | - | - | - | |
| TDS | 1000 | 67 | 91 | |
| ТН | 500 | 65 | 64 | |
| Са | - | - | - | |
| Mg | - | - | - | |
| Na | 200 | 88 | 100 | |
| К | - | - | - | |
| CI | 250 | 73 | 88 | |
| NO ₃ | 50 | 55 | 54 | |
| CO ₃ | - | - | - | |
| HCO₃ | - | - | - | |
| SO ₄ | 250 | 100 | 100 | |

Table 3. Distribution of water samples within the drinking water standards collected from shallow and deep aquifers

All unites have their measurement in mg/l except, 'T' is expressed in °C, 'EC' is expressed in µScm, 'TH and CaH' are expressed in mg/l as CaCO3

Table 4. Water-types for shallow aquifer samples from Naokari and Manikgarh limestone mine area

| Well ID | Pre 2010 | Post 2010 |
|---------|------------------------------|--|
| DW1 | Mg-Ca-Na-HCO₃-Cl | Mg-Ca-HCO ₃ -CO ₃ -Cl |
| DW2 | Mg-Ca- HCO ₃ | Mg-Ca-CI-NO3-HCO ₃ |
| DW3 | Mg-Ca-HCO ₃ -Cl | Mg-Ca-CI-HCO ₃ |
| DW4 | Mg-Ca-HCO ₃ -Cl | Mg-Ca- HCO ₃ -Cl |
| DW6 | Ca-Mg-Na-CI-HCO₃ | Ca-CI-SO ₄ |
| DW7 | Na-Mg-Ca-Cl-HCO₃ | Na-Mg-Ca-HCO₃-Cl |
| DW8 | Mg-Na- HCO₃-Cl | Na-Mg- HCO ₃ -Cl |
| DW9 | Mg-Na-CI- HCO ₃ | Na-Mg- HCO ₃ -Cl |
| DW17 | Mg-Na-Ca-CI-HCO ₃ | Mg-Ca-Na-HCO₃ |
| DW 5 | Ca-Mg- HCO₃-Cl | Mg-Ca-K-SO ₄ - HCO ₃ - NO ₃ |
| DW10 | Mg-Ca- HCO₃ | Mg-Ca-HCO₃ |
| DW11 | Mg-Ca- HCO₃-Cl | Ca-Mg-CI- HCO ₃ -SO ₄ -NO ₃ |
| DW12 | Mg-Ca- HCO ₃ -Cl | Mg-Ca- HCO₃ |
| DW13 | Ca-Mg- HCO₃ | Ca-Mg- HCO₃-Cl |
| DW14 | Mg-Ca- HCO₃-Cl | Ca-Mg- HCO₃-Cl |
| DW15 | Mg-Ca- HCO₃ | Mg-Ca- HCO₃ |
| DW16 | Mg-Ca- HCO₃-Cl | Ca-Mg- HCO ₃ -Cl |

- 1. Alkaline earths exceed alkalis
- 2. Alkalis exceed alkaline earths
- 3. Week acids exceed strong acids
- 4. Strong acids exceed week acids
- 5. Carbonate hardness exceeds 50%
- 6. Non carbonate hardness exceeds 50%
- 7. Non carbonate alkalis exceeds 50%
- 8. Carbonate alkalis exceeds 50%
- 9. No one cation-anion pair exceeds 50%

The water-types revealed through Piper diagram are presented in Table 4 and 5. The data shows that Mg^{2+} is the dominant cation followed by Ca^{2+} and Na^+ . Mg^{2+} is dominant in 64.70% samples and both Ca^{2+} and Na^+ in 35.29% of samples in shallow aquifer. In deep aquifer

samples Mg^{2+} is dominant in 42.30% samples and both Ca^{2+} and Na^+ in 57.70% of samples. HCO_3^- is the most dominant anion (78% samples) followed by CI⁻ and SO₄⁻²⁻. It is observed that in almost all pre-monsoon shallow aquifer samples from Naokari limestone mine area alkaline earths (Ca + Mg) exceed alkalis (Na + K), except one sample from Awarpur village (DW7). 100% pre-monsoon shallow aquifer samples from Naokari limestone mine area have weak acids (CO₃ + HCO₃) exceed the strong acids (SO₄+Cl +F). 25% of samples exhibit that non carbonate hardness (secondary alkalinity) exceeds 50% i.e. chemical properties of the groundwater are dominated by alkalies and strong acids. In pre-monsoon season from Naokari limestone mine

| Well ID | Pre 2010 | Post 2010 |
|---------|---|--------------------------------|
| BW1 | Mg-Na- HCO ₃ | Mg-Na-Ca- HCO₃-Cl |
| BW2 | Na-Mg-Ca- HCO₃ | Na-Mg-Ca- HCO₃-Cl |
| BW3 | Ca-Mg- HCO₃ | Mg-Ca-Na- HCO₃ |
| BW4 | Na-Ca- HCO₃-Cl | Ca-Mg- HCO ₃ |
| BW5 | Mg-Ca- Na-Cl-HCO ₃ | Ca-Na-SO4-HCO ₃ -Cl |
| BW6 | Mg-Ca- HCO₃ | Ca-Mg- HCO₃-Cl |
| BW7 | Na-Mg- HCO ₃ | Mg-Na-HCO ₃ |
| BW8 | Mg-Na-Ca- HCO ₃ | Na-Mg-HCO ₃ |
| BW9 | Ca-Mg-Na- HCO₃-Cl | Ca-Mg-HCO₃-Cl |
| BW11 | Mg- HCO ₃ -SO ₄ ⁻ Cl | Ca-Mg-HCO ₃ |
| BW14 | Na-Mg- HCO ₃ -CO ₃ | Ca-Mg-CI-HCO ₃ |
| BW10 | Ca-Mg-HCO ₃ | Ca-Mg-HCO ₃ |
| BW12 | Mg-Ca- HCO ₃ | Mg-Ca-HCO3-CI |
| BW13 | Mg-Ca- HCO₃ | Ca-Mg-CI-NO3- HCO ₃ |

Table 5. Water-types for Deep aquifer samples from Naokari andManikgarh limestone mine area



| Туре | Chemical Type |
|----------|----------------------------|
| Type I | CaHCO₃ Type |
| Type II | NaCl Type |
| Type III | Mixed CaNaHCO ₃ |
| Type IV | Mixed CaMgCl |
| Type V | CaCl Type |
| Type VI | NaHCO₃ Type |

Figure 2. Characteristic features of Piper Diagram

area 61.11% of deep aquifer samples show alkaline earths (Ca + Mg) exceeds the alkalis (Na + K). About 63.64% of deep samples reveal that weak acids (CO₃ + HCO₃) exceed strong acids (SO₄+Cl +F). Remaining 36.36% samples divulge strong acids (SO₄+Cl +F) exceed weak acids (CO₃ + HCO₃).

Similarly all the pre-monsoon shallow aquifer samples from Manikgarh limestone mine area exhibit that the alkaline earths (Ca + Mg) exceeds alkalis (Na + K) as well as weak acids (CO₃ + HCO₃) exceed strong acids (SO₄+Cl +F). While 73 % deep aquifer samples reveal alkaline earths (Ca + Mg) exceeds alkalis (Na + K) and all the samples represent that the weak acids (CO₃ + HCO₃) exceed strong acids (SO₄+Cl +F).

It is observed from 50% post-monsoon shallow aguifer samples of Naokari limestone mine area that the alkaline earths (Ca + Mg) exceed alkalis while 50% show alkalis (Na + K) exceed alkaline earths (Ca + Mg). 33.33% post-monsoon shallow aquifer samples from Naokari limestone mine area have weak acids (CO₃ + HCO₃) exceed the strong acids (SO₄+Cl+F) while remaining 66.67% samples show strong acids (SO₄+Cl +F) exceed weak acids (CO₃ + HCO₃). 38% of samples exhibit that non carbonate hardness (secondary alkalinity) exceeds 50% i.e. chemical properties of the groundwater are dominated by alkalis and strong acids. Also, 38% of samples exhibit that no one cation-anion exceeds 50%. In post-monsoon season, from Naokari limestone mine area 50.09% of deep aquifer samples show alkaline earths (Ca + Mg) exceed the alkalis (Na+K). About 75% of deep aquifer samples reveal that weak acids $(CO_3 + HCO_3)$ exceed strong acids (SO₄+CI+F). Remaining 50% samples divulge strong acids (SO_4+CI+F) exceed weak acids $(CO_3 + HCO_3)$.

Correspondingly all the post-monsoon shallow aquifer samples from Manikgarh limestone mine area exhibit that the alkaline earths (Ca + Mg) i.e. 100% exceeds alkalis (Na + K) as well as weak acids (CO₃ + HCO₃) i.e. 52.94% exceed strong acids (SO₄+Cl +F) i.e. 47.06%. While 100% deep aquifer well samples reveal alkaline earths (Ca + Mg) exceeds alkalis (Na + K) and also 50% samples represent that the weak acids (CO₃ + HCO₃), strong acids (SO₄+Cl +F) i.e. 50%.

From another point of view 71 % of the plots clustered in Type I (Ca+MgCO3+HCO3) facies of the Piper's diagram. This might be attributed to hydro-geochemistry of the shallow aquifer was mainly controlled by dissolution of carbonate minerals in pre monsoon seasons 2010.

Plots of analytical data of the water samples collected from deep aquifer clustered in the divisions 1,2,3,4 of Piper diagram (Figure 4) in pre-monsoon season, representing equal dominance of alkaline earths and alkali, dominance of weak acid, carbonate hardness and non carbonate alkali. Plots of analytical data of the water samples collected from the same aquifer in post monsoon seasons clustered in the divisions of 1, 4, 5 which indicated equal dominance of alkaline earths and alkali, dominance of strong acids, equal dominance of carbonate hardness and non carbonate hardness and dominance of non-carbonate alkali (Table 4) type of hydro-geochemistry.

Ion exchange process in the aquifers

The ion exchange between the ground water and its host environment during residence or travel can be understood by studying the chloroalkaline indices i.e.

CAI [(CI⁻ Na⁺ +K ⁺)/CI] and CAII[(CI⁻ Na⁺ +K ⁺)/(SO₄²⁻ +HCO₃+CO₃²⁻+NO₃)].

100% of the water samples collected from shallow aquifers and deep aquifers in both post-monsoon and pre-monsoon seasons had positive CA–I and CA–II values. Positive chloroalkaline indices is due to base exchange (Na⁺ and K⁺ ions in water are exchanged with Mg²⁺ and Ca²⁺ ions) reaction (Schoeller, 1965 and 1967).

Mechanism controlling ground water chemistry

Gibbs's diagrams representing the ratios of $[Na + :(Na^{+} +$ Ca²⁺)] and [CI :(CI + HCO₃)] as a function of TDS, are widely employed to assess the functional source of dissolved chemical constituents, such as precipitation dominance, rock dominance and evaporation dominance (Gibbs, 1970). Gibbs's plot (Figure 7) of analytical data of the ground water samples collected from shallow aguifer clustered at the region of rock dominance. This might be attributed to chemical weathering of rock forming minerals is the major driving force in controlling ground water chemistry. Gibbs's diagram plot (Figure 8) of analytical data of the water samples exhibit that chemical weathering of rock forming minerals is the main causative factor in the evolution of chemical composition of groundwater occurring in all the lithological domains of study area, which is later influenced by anthropogenic activities. The ratio of cations Na^+ : $(Na^+ + Ca^+)$ and Cl^- : $(Cl^{-} + HCO_{3})$ are spread from the rock domain towards the zone of precipitation dominance indicating the existence of rock water interactions which is a major source of dissolved ions in the groundwater.

Analysis of irrigation water samples

The electrical conductivity and Na⁺ have a fundamental role in suitability of water for irrigation. According to Saleh et al., (1999) the excess salinity reduces the osmotic activity of plants and hence interferes with the absorption of water and nutrients from the soil. The excess amount of dissolved ions affects the physical and the chemical setup of agricultural soil. In addition, salts affect the



Figure 3. Piper diagram plot of the water samples collected from shallow aquifer (Premonsoon 2010 samples)

(Note: samples from Manikgarh area samples shown by Red colour and Naokari area samples shown by Black colour in piper diagram)



Figure 4. Piper diagram plot of the water samples collected from Deep aquifer (Premonsoon 2010 samples)



Figure 5. Piper diagram plot of the water samples collected from Shallow aquifer (Post-monsoon 2010 samples)



Figure 6. Piper diagram plot of the water samples collected from Deep aquifer (Postmonsoon 2010 samples)



Figure 7. Gibb's diagram plot of the water samples collected from shallow aquifer



Figure 8. Gibb's diagram plot of the water samples collected from deep aquifer

growth of plants directly and also indirectly affect the soil structure, permeability and aeration (Subba Rao, 2006).

Sodium hazards

Values of residual sodium carbonate (RSC), percentage sodium (%Na) and sodium adsorption ratio (SAR) of the collected samples are described in Table 5. RSC is a measurement that compares the concentration of Ca²⁺ and Mg²⁺ to HCO₃ and CO₃². It also determines when Ca²⁺ and Mg²⁺ (macro nutrients) precipitation can occur in the soil and results in additional Na⁺ dominance of soil cation exchange sites.

In the present study, the suitability of groundwater has been checked for irrigation purpose using the Sodium Absorption Ratio (SAR) {SAR = $Na^+ / \sqrt{(Ca^{++} +)}$ Mg⁺⁺)/2]} and the US Salinity Laboratory diagram (1954) (Figure 9.1 to 9.4). In addition, the Percent Sodium (%Na), Residual Sodium Carbonate (RSC), Residual Sodium Bicarbonate (RSBC), Soluble Sodium Percentage (SSP), Corrosively Ratio (CR), Kelley's Ratio (KR) and permeability Index (PI) have been also estimated which decides its suitability for irrigation purpose. Thus, suitability of groundwater from the study area for irrigation purpose is calculated and the data are presented in Table 6 and 7).

Sodium Absorption Ratio (SAR)

A high sodium concentration changes soil properties and reduce soil permeability. The suitability of water for irrigation purpose is determined by the mathematical equation:

SAR= Na/ $\sqrt{Ca+Mg/2}$. (Where, ions are expressed as meq/l.)

Water with SAR value ≤ 10 is considered as of excellent quality, 10 to 18 is good, between 18 to 26 is fair and above 26 is said to be unsuitable for irrigation. The calculated SAR values from Naokari and Manikgarh limestone mine areas are less than 10 meg/l [pre monsoon 2010 (shallow aguifer: 0.116 to 1.586; deep aguifer: 1.198 to 1.875) post monsoon 2010 (deep aquifer: 0.103 to 1.90; Bore well: 0.121 to 1.226). In general, the plots of the groundwater samples from both from Naokari and Manikgarh limestone mine areas fall in C_2 - S_1 and C_3 - S_1 category, which indicate the water of medium to high salinity-low sodium type and can be utilized for irrigation with very little danger of exchangeable sodium (Murkute, 2014).

Residual Sodium Carbonate (RSC)

RSC value (in meq/l) is calculated as RSC = $(HCO_3^{-} + CO_3^{-}) - (Ca^{++} + Mg^{++})$. A high RSC value in water leads to an increase in the adsorption of sodium on soil (Eaton 1950). RSC level less than 1.25meq/l is considered safe,

where as water with RSC of 1.25 - 2.50 meq/l is within marginal range and RSC value of water sample 2.50 meq/l or greater is considered too high making the water unsuitable for irrigation use. The maximum RSC value of groundwater sample is noted from pre-monsoon deep aquifer sample, collected from BW2- Asan kh village located in Naokari limestone mine area. In pre-monsoon 2010, shallow aquifers having 10.782 to -0.317 and deep aquifer ranging -7.544 to 2.052, in post-monsoon 2010, shallow aquifer is -16.663 to 1.704 and deep aquifer ranging from 24.341 to 0.729. RSC value negative which may be attributed to excess Ca²⁺ and Mg²⁺ in irrigation water.

Percent Sodium (%Na)

The % Na is widely used for evaluating the suitability of water quality for irrigation (Wilcox, 1955). This value is defined by the expression: % Na = Na + K/ (Ca + Mg + Na + K) x 100. High % Na in irrigation water causes exchange of sodium in water, and exchange of calcium and magnesium contents in soil having poor internal drainage. The % Na, < 60 represents safe water while it is unsafe if > 60 (Eaton, 1950). The% Na values from Naokari and Manikgarh limestone mine areas do not exceed the limit of 60 as stated by Eaton (1950), except the one bore well sample (BW4- Village Sonurli from Naokari limestone mine area. [Pre-monsoon 2010 (shallow aquifer: 7.95 to 52.39; deep aquifer: 12.09 to 65.30) post-monsoon 2010 (shallow aquifer: 6.62 to 61.76; deep aquifer: 9.43 to 47.96). As per this criterion the groundwater from Naokari and Manikgarh limestone mine areas is safe for irrigation purpose.

ΡI

In the study area, PI values fall in class I of Doneen (1964) sample from pre-monsoon 2010 shallow aquifer in between 48.749 to 110.69 and deep aquifer 31.761 to 86.562 (Table 4, 5). In post-monsoon 2010, shallow aquifer shows 17.827 to 77.051 and deep aquifer are in between 33.173 to 85.138 (Table 6, 7). On the basis of computed PI values, it can be understood that the groundwater from Naokari and Manikgarh limestone mine areas is suitability for the irrigation purpose.

Ca ²⁺ /Mg ²⁺ ratio

 Ca^{2+}/Mg^{2+} ratio of the 57.14% of collected irrigation water samples in post-monsoon and 71.43% of collected irrigation water samples in pre-monsoon were less than one. It indicated that, these were Mg ²⁺ dominated water. In Mg²⁺ dominated water the potential effect of Na⁺ may be slightly increased (FAO, 1994) might be due to Mg²⁺ induced Ca²⁺ deficiency caused by high levels of exchangeable Mg²⁺ in soil. Ca²⁺ appears to reduce



Figure 9.1. US Salinity Diagram of groundwater samples from Naokari and Manikgarh limestone mine areas (Pre-monsoon 2010 Shallow aquifer)



Figure 9.2. US Salinity Diagram of groundwater samples from Naokari and Manikgarh limestone mine areas (Pre-monsoon 2010 Deep aquifer)



Figure 9.3. US Salinity Diagram of groundwater samples from Naokari and Manikgarh limestone mine areas (Post-monsoon 2010 Shallow aquifer)



Figure 9.4. US Salinity Diagram of groundwater samples from Naokari and Manikgarh limestone mine areas (Post monsoon 2010 Deep aquifer)

| Sample No | | SAR | % Na | RSC | RSBC | CR | SSP | KR | PI |
|-----------|-----------|-------|-------|---------|---------|-------|--------|-------|--------|
| 1 | Pre 2010 | 0.367 | 23.36 | -1.385 | 2.809 | 0.185 | 23.369 | 0.276 | 88.078 |
| | Post 2010 | 0.250 | 16.86 | -1.592 | -0.0223 | 0.079 | 16.860 | 0.198 | 37.866 |
| 2 | Pre 2010 | 0.227 | 16.89 | -1.653 | 1.060 | 0.064 | 16.899 | 0.198 | 80.162 |
| | Post 2010 | 0.233 | 9.22 | -16.663 | -4.978 | 0.680 | 9.226 | 0.098 | 17.827 |
| 3 | Pre 2010 | 0.190 | 11.25 | -4.412 | 2.004 | 0.247 | 11.257 | 0.124 | 79.552 |
| | Post 2010 | 0.309 | 14.05 | -10.849 | -3.030 | 0.467 | 14.055 | 0.156 | 24.140 |
| 4 | Pre 2010 | 0.248 | 14.87 | -2.130 | 3.134 | 0.251 | 14.879 | 0.171 | 85.245 |
| | Post 2010 | 0.330 | 15.02 | -8.832 | 1.0386 | 0.584 | 15.052 | 0.170 | 28.717 |
| 5 | Pre 2010 | 0.216 | 15.51 | -1.365 | 1.513 | 0.106 | 15.517 | 0.179 | 77.521 |
| | Post 2010 | 0.464 | 38.7 | -3.609 | 1.7372 | 0.370 | 38.708 | 0.296 | 42.424 |
| 6 | Pre 2010 | 0.527 | 25.73 | -10.782 | -3.872 | 0.709 | 25.738 | 0.271 | 50.006 |
| | Post 2010 | 0.375 | 15.56 | -14.122 | -11.242 | 0.483 | 15.569 | 0.18 | 24.072 |
| 7 | Pre 2010 | 1.586 | 52.39 | -1.237 | 3.369 | 1.105 | 52.394 | 1.091 | 90.574 |
| | Post 2010 | 0.852 | 40.46 | 0.255 | 2.809 | 0.260 | 40.462 | 0.667 | 62.025 |
| 8 | Pre 2010 | 0.745 | 37.18 | -0.534 | 4.153 | 0.226 | 37.188 | 0.588 | 110.69 |
| | Post 2010 | 1.538 | 48.9 | -2.227 | 7.032 | 1.469 | 48.907 | 0.853 | 58.606 |
| 9 | Pre 2010 | 1.265 | 46.5 | -4.799 | 4.989 | 1.752 | 46.501 | 0.681 | 92.002 |
| | Post 2010 | 1.900 | 61.76 | 1.704 | 6.146 | 0.835 | 61.764 | 1.470 | 77.051 |
| 10 | Pre 2010 | 0.203 | 14.76 | -0.497 | 2.628 | 0.096 | 14.766 | 0.164 | 83.467 |
| | Post 2010 | 0.159 | 12.15 | -0.981 | 2.802 | 0.072 | 12.153 | 0.129 | 44.474 |
| 11 | Pre 2010 | 0.180 | 16.28 | -2.0585 | 0.327 | 0.067 | 16.289 | 0.167 | 76.008 |
| | Post 2010 | 0.168 | 12.59 | -8.480 | -3.544 | 0.495 | 12.592 | 0.093 | 23.519 |
| 12 | Pre 2010 | 0.140 | 9.45 | -3.090 | 2.010 | 0.148 | 9.4532 | 0.101 | 83.775 |
| | Post 2010 | 0.134 | 10.71 | -2.059 | 1.3956 | 0.087 | 10.719 | 0.104 | 38.644 |
| 13 | Pre 2010 | 0.147 | 12.14 | -0.317 | 1.245 | 0.071 | 12.142 | 0.132 | 68.531 |
| | Post 2010 | 0.165 | 12.26 | -3.720 | -0.101 | 0.148 | 12.263 | 0.115 | 33.443 |
| 14 | Pre 2010 | 0.257 | 14 | -4.994 | 0.270 | 0.364 | 14.007 | 0.160 | 58.410 |
| | Post 2010 | 0.272 | 14.82 | -5.444 | -0.344 | 0.349 | 14.823 | 0.166 | 32.661 |
| 15 | Pre 2010 | 0.192 | 15.63 | -0.861 | 1.441 | 0.045 | 15.632 | 0.179 | 86.125 |
| | Post 2010 | 0.280 | 16.63 | -4.076 | 0.4475 | 0.165 | 16.638 | 0.190 | 36.779 |
| 16 | Pre 2010 | 0.116 | 7.95 | -4.437 | -0.982 | 0.126 | 7.9546 | 0.083 | 48.749 |
| | Post 2010 | 0.103 | 6.62 | -7.581 | -2.316 | 0.210 | 6.624 | 0.061 | 21.860 |
| 17 | Pre 2010 | 0.676 | 31.21 | -4.207 | 0.974 | 0.927 | 31.211 | 0.448 | 77.728 |
| | Post 2010 | 1.00 | 41.41 | -2.626 | 0.663 | 0.355 | 41.410 | 0.694 | 57.904 |

Table 6. Irrigation quality results of groundwater samples (shallow aquifer 2010)

possible toxicities due to other ions like Na⁺, Mg²⁺ in the root environment. If the Ca²⁺ /Mg²⁺ ratio is near or less than one, the uptake and translocation of Ca²⁺ from soil water to the above ground parts of the growing crop is diminished due to antagonistic effects of high Mg²⁺ or competition for absorption sites to such an extent that less Ca²⁺ is absorbed.

Concentration of common phytotoxic anions

Values of concentration of common phytotoxic anions (Cl and F) of the collected samples are described in Table 5. In the irrigation water the most common toxicity is from Cl. Cl is not adsorbed or held back by soils, therefore it moves readily with the soil water, Is taken up by crops, moves in the transpiration stream, and accumulates in

leaves. If the CI concentration in the leaves exceeds the tolerance of the crops, injury symptoms develop such as leaf burn or drying of leaf tissue. With sensitive crops, these symptoms occur when leaves accumulate from 0.3 to 1.0% CI on a dry weight basis, but sensitivity varies among these crops. Usual range of chloride in irrigation water is 0-30 meq/l (FAO, 1994). In the collected irrigation water samples CI concentration ranged between 28-438 and 28-363 in pre and post monsoon seasons respectively. 22% of the irrigation water samples collected in post-monsoon and 27% of the irrigation water samples collected in pre-monsoon were found within the usual range.

F⁻ in the soil is generally not harmful. But F⁻ on plant surfaces may be harmful to plants and grazing animals. BIS (1991) recommended 1mg/l maximum

| Sample No | | SAR | % Na | RSC | RSBC | CR | SSP | KR | PI |
|-----------|-----------|-------|-------|---------|--------|-------|--------|-------|----------|
| 1 | Pre 2010 | 0.511 | 26.3 | -1.016 | 5.646 | 0.227 | 26.302 | 0.353 | 50.07877 |
| | Post 2010 | 0.563 | 25.19 | -5.065 | 10.458 | 0.441 | 25.193 | 0.324 | 49.479 |
| 2 | Pre 2010 | 1.338 | 54.73 | 2.052 | 4.685 | 0.190 | 54.736 | 1.194 | 78.54496 |
| | Post 2010 | 1.141 | 40.72 | -4.0385 | 11.371 | 0.637 | 40.721 | 0.684 | 63.745 |
| 3 | Pre 2010 | 0.213 | 16.71 | -1.528 | 1.515 | 0.123 | 16.717 | 0.160 | 42.58443 |
| | Post 2010 | 0.358 | 22.08 | -2.261 | 8.344 | 0.118 | 22.087 | 0.269 | 54.488 |
| 4 | Pre 2010 | 1.875 | 65.3 | 0.498 | 1.978 | 0.292 | 65.303 | 1.869 | 83.56262 |
| | Post 2010 | 0.121 | 9.43 | -2.659 | 10.028 | 0.111 | 9.4305 | 0.090 | 43.833 |
| 5 | Pre 2010 | 0.456 | 25.61 | -7.544 | -0.716 | 0.537 | 25.615 | 0.262 | 34.74854 |
| | Post 2010 | 0.551 | 23.38 | -24.341 | 22.929 | 2.266 | 23.387 | 0.199 | 33.173 |
| 6 | Pre 2010 | 0.217 | 13.31 | -2.344 | 2.837 | 0.184 | 13.319 | 0.150 | 38.57248 |
| | Post 2010 | 0.246 | 15.69 | -4.110 | 10.880 | 0.288 | 15.694 | 0.159 | 43.797 |
| 7 | Pre 2010 | 1.104 | 44.77 | 0.377 | 5.971 | 0.310 | 44.772 | 0.807 | 65.38937 |
| | Post 2010 | 0.852 | 37.54 | 0.0230 | 10.871 | 0.334 | 37.548 | 0.592 | 64.690 |
| 8 | Pre 2010 | 0.818 | 38.5 | 1.078 | 5.603 | 0.108 | 38.501 | 0.622 | 63.55452 |
| | Post 2010 | 1.226 | 46.91 | 0.729 | 8.867 | 0.220 | 46.914 | 0.880 | 73.004 |
| 9 | Pre 2010 | 0.627 | 27.96 | -3.585 | 0.527 | 0.624 | 27.962 | 0.385 | 45.86813 |
| | Post 2010 | 0.381 | 18.32 | -7.650 | 14.027 | 0.320 | 18.328 | 0.212 | 42.816 |
| 10 | Pre 2010 | 0.234 | 19.18 | -0.871 | 1.020 | 0.065 | 19.186 | 0.211 | 51.17352 |
| | Post 2010 | 0.372 | 15.46 | -12.763 | 16.119 | 0.236 | 15.469 | 0.180 | 37.481 |
| 11 | Pre 2010 | 0.206 | 12.09 | -4.899 | -1.280 | 0.182 | 12.090 | 0.134 | 31.76172 |
| | Post 2010 | 0.152 | 9.96 | -4.1634 | 10.773 | 0.172 | 9.9623 | 0.102 | 41.103 |
| 12 | Pre 2010 | 0.198 | 17.31 | -0.904 | 1.562 | 0.053 | 17.311 | 0.191 | 51.99549 |
| | Post 2010 | 0.826 | 47.96 | 0.6282 | 3.171 | 0.041 | 47.960 | 0.871 | 85.138 |
| 13 | Pre 2010 | 0.215 | 22.49 | -1.319 | 2.628 | 0.079 | 22.493 | 0.164 | 43.51718 |
| | Post 2010 | 0.257 | 22.8 | -11.517 | 13.914 | 0.548 | 22.807 | 0.126 | 34.427 |
| 14 | Pre 2010 | 0.826 | 47.96 | -0.970 | 2.073 | 0.041 | 47.960 | 0.871 | 70.65757 |
| | Post 2010 | 0.220 | 12.08 | -5.556 | 8.0283 | 0.176 | 12.088 | 0.135 | 40.696 |

Table 7. Irrigation quality results of groundwater samples from (deep aquifer 2010)

concentrations of F⁻ in irrigation water. F⁻ concentration in the collected from shallow aquifer in pre-monsoon seasons water samples ranged between 0.53-7.52 mg/l and samples collected from deep aquifer in postmonsoon seasons range from 0.48-6.28 mg/l. This implied the irrigation waters are not safe from F point of view.

CONCLUSION

Quality of the water samples collected from the deep aquifer found satisfactory in comparison with water samples collected from the shallow aquifer. It was concluded from the Piper diagram plot that hydrogeochemistry of confined aquifer was earth alkaline with increased portion of alkalis and with prevailing bicarbonate. Chloro-alkaline disequilibrium type of ion exchange process was found in shallow aquifer. Chemical weathering of rock forming minerals was the major driving force controlling ground water chemistry as represented by Gibbs diagram plot. Hydro-geochemistry of deep aquifer in post monsoon season was alkaline with bicarbonate and sulphate chloride and earth alkaline with increased portion of alkalis with prevailing bicarbonate where as in pre-monsoon season. Hydrogeochemistry of the deep aquifer shifted towards alkaline with sulphate and bicarbonate. Ion exchange process in deep aquifer was found Base Exchange type in both the seasons. According to Gibbs's diagram plot chemical weathering of rock forming minerals in post monsoon season and chemical weathering along with evaporation in pre monsoon season were the major driving force controlling hydro-geochemistry of deep aquifer. RSC, %Na, SAR, Ca²⁺/Mg²⁺ ratio, concentration of CI, F in irrigation water along with US Salinity Laboratory's and Wilcox's diagrams suggested that the majority of water samples were not good for irrigation during pre monsoon season in comparison to post monsoon season.

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